

DURATION OF ACTIVITY ON LOBATE-SCARP THRUST FAULTS ON MERCURY. Maria E. Banks^{1, 2}, Nadine Barlow³, Christian Klimczak⁴, Zhiyong Xiao^{5, 6}, Thomas R. Watters², Clark R. Chapman⁷.

¹Planetary Science Institute, Tucson, AZ, USA, banks@psi.edu. ²Center for Earth and Planetary Studies, National Air and Space Museum, Smithsonian Institution, Washington, DC 20560, USA. ³Northern Arizona University, Flagstaff, AZ, USA. ⁴Department of Geology, University of Georgia, Athens, GA 30602, USA. ⁵School of Earth Sciences, China University of Geosciences (Wuhan), Wuhan, Hubei, 430075, China. ⁶Centre for Earth Evolution and Dynamics, University of Oslo, Sem Sælands vei 2A, 0371 Oslo, Norway. ⁷Department of Space Studies, Southwest Research Institute, Boulder, CO 80302, USA.

Introduction: Contractual deformation on Mercury is expressed by three main types of landforms: lobate scarps, high-relief ridges, and wrinkle ridges [e.g., 1–4], with lobate scarps (rupēs) the most widely distributed of these landforms. Lobate scarps deform all major geologic units, including intercrater plains and smooth plains, thereby providing valuable insight into the history of horizontal shortening on Mercury. Lobate scarps are interpreted to be the expression of low-angle (<45° fault dip) surface-breaking thrust faults, and they can extend more than 500 km in length and display up to ~3 km of relief [1, 3, 5]. Their formation has been attributed primarily to compressional stresses produced by planetary cooling and global contraction [1–3, 5]. Understanding the history of crustal deformation provides constraints on thermal history models and insight into the interplay between tectonics and volcanism and the cooling and solidification of the planet’s interior [5].

Methods and data: The recent crater production function and inner solar system chronology of [6] indicates that the oldest surfaces on Mercury date from about 4.0–4.1 Ga, during the Late Heavy Bombardment (LHB), and correspond approximately to the pre-Tolstojan and Tolstojan systems [7]. Widespread smooth volcanic plains were emplaced by about 3.55–3.8 Ga [8], at the end of the Calorian system [7]. The Calorian is followed chronologically by the Mansurian and Kuiperian systems [7]. Craters of different ages exhibit different amounts of degradation ranging from sharp morphologies and the presence or absence of high reflectance ejecta and rays (Kuiperian and Mansurian craters, respectively), to moderately degraded (Calorian craters), and heavily degraded (Tolstojan and pre-Tolstojan craters) morphologies, characterized by subdued rims, along with infilling of the crater floor, and superposing craters [7, 9].

Orbital images and mosaics from the Mercury Surface, Space ENvironment, GEochemistry, and Ranging (MESSENGER) Mercury Dual Imaging System (MDIS) are used to investigate previously unrecognized stratigraphic relationships between lobate scarps, and impact craters exhibiting a range of degradation states, to gain further insight into the duration of thrust fault activity on Mercury [9]. Here we focus on Mercury’s 30 officially named scarps (Fig. 1). For 22 of these

lobate scarps, size–frequency distributions were also determined for craters that intersect the scarp face and immediately surrounding surface. This was completed using a modified buffered crater counting technique [e.g., 10–11], and only craters for which the centers overlapped the count area were included. Crater count-derived ages were obtained using the model production function chronology of Mercury from [6]. Age estimates for impact craters that are crosscut by scarp segments provide lower limits on the age of the most recent activity on the underlying thrust faults, but do not constrain the time of formation of those faults. Age estimates for craters that superpose scarp segments provide upper bounds on the age of the most recent detectable activity on the thrust faults associated with the scarp segments, and constrain the time of formation of the fault and the initiation of slip to before the crater-forming impacts [9].

Duration of lobate scarp activity: *Constraints on the earliest thrust faulting.* Any record of crustal deformation prior to or in the early stages of the LHB is unlikely to be preserved. Also, no evidence of embayment of lobate scarps by early-emplaced smooth plains material has been found [4]. Preliminary results from counts of craters crosscut by scarp segments indicate that the majority of the investigated structures formed on surfaces that are Calorian and date back to near the end of the LHB. All of the named scarps collocated with Tolstojan and pre-Tolstojan craters, crosscut and deform the craters (Fig. 2), and all but 7 of the named scarps crosscut Calorian craters. The oldest craters observed to superpose the scarps in this study, or any other scarp segments on Mercury so far, are estimated to have formed during the Calorian system and constrain the time of formation of the associated fault and the initiation of slip to before the end of this system (Fig. 2) [9]. For scarps where a sufficient number of superposing craters could be measured (14 of the investigated scarps), preliminary age estimates also suggest that observable activity may have ceased on some scarp segments before the end of the Calorian.

These collective observations support initiation of shortening of Mercury’s surface on at least a regional scale by some time before the end of the Calorian (before ~3.6 Ga), a time interval during which the major expanses of smooth plains were emplaced. Although

Figure 3. Mercury's named scarps categorized by estimated upper and lower limits for the age of most recent observable activity on the underlying thrust faults.